

NATURAL AND FORCED FLOWS INVOLVING BUOYANCY

- fluid flows which occur in a natural environment or industrial setting and involve the interaction of fluids of different density. Density contrasts may arise from differences in temperature, concentration or a particular quantity.... - High Reynolds # flows (plumes, gravity currents) - Low Reynolds # flows (gravity driven flows and plumes in porous media).

Turbulent, high Re plumes: $Re = \frac{\rho_w b}{\eta} > 10^3 - 10^4$

+ plume is a vertical buoyancy driven flow, so named because of a conical time averaged geometry and bulbous head (if finite height).

consider an axisymmetric plume emanating from $z=0$ consisting

of isotropically mixed, fully turbulent fluid of density $\rho(z)$. let external/ambient fluid outside the time averaged plume envelope $b(z)$ (radius) be of density $\rho_a(z)$ and be radially entrained into plume at $b(z)$ at velocity $u_E(z)$. let plume velocity field (time averaged) be $u(r, z)$.

$$\text{or } 0 < r \leq b(z).$$

relate plume density to ambient and difference in some conserved quantity C . (e.g. salt, particles, bubbles, heat....)

define useful quantities: Flowrate $Q = \int_0^{b(z)} u(r, z) \cdot 2\pi r dr$

$$\text{buoyancy flux } B = g' Q \quad (4)$$

assert 'separable' velocity field $u(r, z) = u(r) w(z)$ (5). (i.e. $u(r)$ Gaussian Top hat $u(r) = \begin{cases} 1 & r \leq b \\ 0 & r > b \end{cases}$)

to characterise the plume would like to find explicit z dependence of $w(z)$, $b(z)$. USE CONSERVATION LAWS to find this. consider thin section of plume illustrated above.

$$\text{MASS CONSERVATION} \quad \int_0^{b(z)} \rho(z) u(r, z) \cdot 2\pi r dr = \int_{z=0}^z u_E(z) \rho_a(z) \cdot 2\pi b(z) dz + \rho_a \rho_0 \quad (6)$$

$$\text{MOMENTUM CONSERVATION} \quad d \left(\int_0^{b(z)} \rho(z) u^2(r, z) \cdot 2\pi r dr \right) = (\rho_a - \rho) \cdot \pi b^2 g dz \quad (7)$$

rate of change of momentum = imbalance of Archimedian upthrust and gravitational force.

$$\text{"C" CONSERVATION} \quad \int_0^b \rho(z) u(r, z) C(z) \cdot 2\pi r dr = \int_{z=0}^z u_E(z) \rho_a(z) C_a(z) \cdot 2\pi b(z) dz + C_a \rho_0 Q_0 \quad (8)$$

For closure hypothesis relate entrainment velocity $u_E(z)$ to plume velocity at $b(z) = r$. "C" same for

$$i.e. u_E(z) = \epsilon u(b, z) \quad (t). \quad \epsilon \text{ is entrainment coefficient. (Typically } \sim 0.1).$$

using a Top Hat velocity field i.e. $u(r, z) = \begin{cases} w(z) & r \leq b(z) \\ 0 & r > b(z) \end{cases}$, conservation laws become: (in differential form)

$$\rightarrow \text{MASS} \quad \frac{d}{dz} (\rho_w b^2) = 2 \epsilon w b \rho_a \quad (9)$$

$$\text{MOM.} \quad \frac{d}{dz} (w^2 b^2 \rho) = \rho_a \beta (c_a - c) b^2 g \quad (10)$$

$$\text{"C".} \quad \frac{d}{dz} (\rho_w b^2 C) = 2 \epsilon w b \rho_a c_a \quad (11)$$

$$\text{Now from (9)} \quad \frac{d}{dz} (\rho_w b^2) = 2 \epsilon w b \rho_a c_a. \quad (12)$$

$$\text{From (12)} \Rightarrow \frac{d}{dz} (\rho_w b^2 [c_a - c]) = \rho_w b^2 \frac{dc_a}{dz} \quad (13)$$

$$\text{Now using (11, 13); } B = -g \beta (c_a - c) Q. \quad \text{using top hat velocity field: } Q = w \pi b^2 \quad (14)$$

$$\therefore B = -\pi g \beta (c_a - c) w b^2 \quad (15).$$

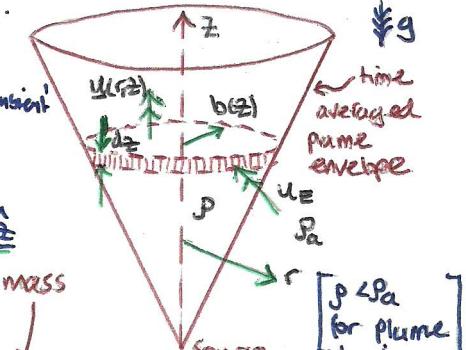
$\rho_a \rho_0$ in all terms apart from those involving $\rho - \rho_a$ constant (i.e. (1) is used) $\Rightarrow (9), (10), (13)$ become

$$\frac{d}{dz} (w b^2) = 2 \epsilon w b \quad (16)$$

$$\frac{d}{dz} (w^2 b^2) = \beta g (c_a - c) b^2 \quad (17)$$

$$\frac{d}{dz} (w b^2 c_a) = w b^2 \frac{dc_a}{dz} \quad (18)$$

[we will consider axisymmetric plumes, deviation is analogous for line plumes, though results different]



$$g' = g \left(\frac{\rho - \rho_a}{\rho_a} \right) \quad (3)$$

$$\text{Gaussian } u(r) = e^{-\frac{r^2}{2b^2}} \quad \text{Top hat } u(r) = \begin{cases} 1 & r \leq b \\ 0 & r > b \end{cases}$$

"Flux of mass is balanced by fluid entrained at boundary b at velocity u_E "

Note closure hypothesis (t) may be extended for certain "Non-Boussinesq" plumes to $u_E = \epsilon w(z) (\rho / \rho_a)^{\delta}$. (12)

$$\text{Consider } \frac{d}{dz} (\rho_w b^2 [c_a - c]) = - \frac{d}{dz} (\rho_w b^2 c_a) + \frac{d}{dz} (\rho_w b^2 c_a) = -2 \epsilon w b \rho_a c_a + c_a \frac{d}{dz} (\rho_w b^2) + \rho_w b^2 \frac{dc_a}{dz}.$$

Application of "Boussinesq approximation", that is: (i.e. (1) is used) $\Rightarrow (9), (10), (13)$ become

(16) - (18) can be recast in terms of variables $Q = \pi w b^2$, $M = w^2 b^2 \pi$ and $B = \pi g \beta w b^2 (c_a - c)$.

(16): $\frac{dQ}{dz} = 2\pi \varepsilon M^{1/2}$ (17): $\frac{dM}{dz} = -\frac{1}{\pi^2} \frac{BQ}{M}$ (18): $\frac{dB}{dz} = -Q g \beta \frac{dc_a}{dz}$ let $N^2 = g \beta \frac{dc_a}{dz}$
 $(N = \text{Brunt-Väisälä frequency})$

∴ plume equations are: (I) $\frac{dQ}{dz} = 2\varepsilon M^{1/2}$

(II) $\frac{dM}{dz} = -\frac{BQ}{M}$ (III) $\frac{dB}{dz} = -N^2 Q$